The Tibetan Plateau-Rocky Mountains Circumglobal (TRC) Wave Train and S2S Precipitation Prediction Regionally and Globally: Results from the GEWEX/LS4P Project

> Xue, Y, A. Bonne, T. Yao, I. Diallo, X. Zeng,, F. Vitart, Q. Tang, D. Neelin, W. K.-M. Lau, LS4P Team.

> > The Weeks3-4/S2S Webinar July 1 2024

Impact of Initialized Land Temperature and Snowpack on Sub-Seasonal to Seasonal Prediction (LS4P)

Inception in the 2018 GEWEX Science Conference Alberta, Canada, 6–11 May 2018

Co-Chairs: Yongkang Xue, Aaron Boones, Tandong Yao

GEWEX/LS4P Project Goals:

What is the impact of the initialization of large-scale land surface temperature/subsurface temperature (LST/SUBT) in high mountains in climate models on the S2S prediction over different regions?
What is the relative role and uncertainties in these land processes versus in SST in S2S prediction?

Xue et al., (2021, GMD)



Observational Evidence T2m Difference be warm and cold Mays



Observed monthly T2m difference between warm and cold years (°C). (a) over the TP based on CMA data; (b) over the western US (WUS) based on NARR data. Note: Years are selected based on TP and WUS May T2m anomalies, using a threshold of 0.5 standard deviation during the period 1981-2010. The years are applied for all months.

Observed difference of surface & subsurface temperature and snow between year with warmest and coldest springs



Liu et al., 2020, JGR

Cold: 1982, 1983, 1986, 1990, 2001 **Warm**: 1999, 2004, 2007, 2009, 2010, 2015

(a) First May 2-m temperature MCA Mode



(a) First Way 2-in temperature with A would



 $\Delta 2_{-\sigma}, (\Delta, \Omega^2 - \sigma, \sigma \Delta^2) = 0.05 (\Omega^1 - \sigma, \sigma^2) (\Omega^1 - \sigma, \sigma^2) (\Omega^2 - \sigma, \sigma^$



First MCA Mode (May 2-m Temperature vis June Precipitation) MCA: Maximum Covariance Analysis (Xue et al., 2018)

Phase I: Tibetan Plateau LST/SUBT Effect is the focus; June 2003 is the first case.









Observed 2 meter May temperature of

Tibetan Plateau over 4000 m

Obs. May 2003 T2m Anomaly (°C)

LS4P Ensemble mean May 2003 T2m

Bias (°C)

Development of Initialization Methodology

Principles Consideration: (1) Model bias; (2) Observed Anomalies; (3) Tuning parameter

Applying the mask, $\tilde{T}_0(i, j)$, will be defined as follows:

 $\tilde{T}_{0}(i, j) = T_{0}(i, j) + \Delta T_{\text{mask}}(i, j) = T_{0}(i, j) + \left[-n \times T_{\text{obs anomaly}}(i, j) - T_{\text{bias}}(i, j)\right],$ when $\bar{T}_{\text{obs anomaly}} \times \bar{T}_{\text{bias}} \ge 0,$

Xue et al. (2021, GMD)

-2 -1 -0.5 0.5 1

2 3 4 5

-5

-3

Simulated May T-2m difference after imposing mask for initial condition

Tibetan Plateau – Rocky Mountain Circumglobal Wave Train (TRC)and TP LST/SUBT Effect Hotspots

0.5 -

The schematic demonstrates the TRC global influence and possible hotspots. The color shadings within the boxes are snapshots of the LS4P multi-model--simulated June 2003 precipitation anomalies due to the effect of cold Tibetan Plateau land surface and subsurface temperature (LST/SUBT), and elsewhere the shaded areas show the observed 200-hPa geopotential height (GHT) anomalies due to the cold Tibetan Plateau temperature. The green bar corresponds to the observations and the red bar is the ensemble mean in each hot spot. Green dots represent a statistical significance at p<0.1. The light vectors are wave activity flux, and the heavy blue arrows indicate the TRC propagation. The figure is based on Xue et al. (BAMS, 2022, Climate Dynamics 2023).

Comparison of June 2003 Precipitation Anomaly due to LST/SUBT and SST Effect(mm/day)

TP LST/SUBT hot spots

Global SST hot spots

TRC Wave Train Effects

n) differences due to the TP **Figure 9.** Ensemble mean of May geopotential height (r. we train simulations and (b) LST/SUBT effect (a) for the 5 models with best TRC was for the 5 models with relatively poor TRC wave train simulations.

Zhang Y. et al., 2024, Climate Dynamics

Figure 5. Non-zonal geopotential height at 200 hPa (m) from (a) ERAI, (b) CIESM EXP0-Nudg, (c) CIESM EXP0, (d) E3SMv1 EXP0-Nudg, and (e) E3SMv1 EXP0 on April 30th, 2003.

Qin Y, Q. Tang et al., Climate Dynamics 2023

Publications

Geosci. Model Dev., 14, 4465–4494, 2021 https://doi.org/10.5194/gmd-14-4465-2021 © Author(s) 2021. This work is distributed under the Creative Commons Attribution 4.0 License.

Impact of Initialized Land Surface Temperature and Snowpack on Subseasonal to Seasonal Prediction Project, Phase I (LS4P-I): organization and experimental design

Yongkang Xue¹, Tandong Yao², Aaron A. Boone³, Ismaila Diallo¹, Ye Liu¹, Xubin Zeng⁴, William K. M. Lau⁵, Shiori Sugimoto⁶, Qi Tang⁷, Xiaoduo Pan², Peter J. van Oevelen⁸, Daniel Klocke⁹, Myung-Seo Koo¹⁰, Tomonori Sato¹¹, Zhaohui Lin¹², Yuhei Takaya¹³, Constantin Ardilouze³, Stefano Materia¹⁴, Subodh K. Saha¹⁵, Retish Senan¹⁶, Tetsu Nakamura¹¹, Hailan Wang¹⁷, Jing Yang¹⁸, Hongliang Zhang¹⁹, Mei Zhao²⁰, Xin-Zhong Liang⁵, J. David Neelin¹, Frederic Vitart¹⁶, Xin Li², Ping Zhao²¹, Chunxiang Shi²², Weidong Guo²³, Jianping Tang²³, Miao Yu²⁴, Yun Qian²⁵, Samuel S. P. Shen²⁶, Yang Zhang²³, Kun Yang²⁷, Ruby Leung²⁵, Yuan Qiu¹², Daniek Peano¹⁴, Xin Qi¹⁸, Yanling Zhan¹², Michael A. Brunke⁴, Sin Chan Chou²⁸, Michael Ek²⁹, Tianyi Fan^{18,10}, Hong Guan³⁰, Hai Lin³¹, Shunlin Liang³², Helin Wei¹⁷, Shaocheng Xie⁷, Haoran Xu⁵, Weiping Li³³, Xueli Shi³³, Paulo Nobre²⁸, Yan Pan²³, Yi Qin^{27,7}, Jeff Dozier³⁴, Craig R. Ferguson³⁵, Gianpaolo Balsamo¹⁶, Qing Bao³⁶, Jinming Feng¹², Jinkyu Hong³⁷, Songyou Hong¹⁰, Huilin Huang¹, Duoying Ji¹⁸, Zhenming Ji³⁸, Shichang Kang^{39,40}, Yanluan Lin⁷⁷, Weiguang Liu^{41,24}, Ryan Muncaster³¹, Patricia de Rosnay¹⁶, Hiroshi G. Takahashi⁴², Guiling Wang⁴¹,

2022

Spring Land Temperature in Tibetan Plateau and Global-Scale Summer Precipitation

Initialization and Improved Prediction

Yongkang Xue, Ismaila Diallo, Aaron A. Boone, Tandong Yao, Yang Zhang, Xubin Zeng, J. David Neelin, William K. M. Lau, Yan Pan, Ye Liu, Xiaoduo Pan, Qi Tang, Peter J. van Oevelen, Tomonori Sato, Myung-Seo Koo, Stefano Materia, Chunxiang Shi, Jing Yang, Constantin Ardilouze, Zhaohui Lin, Xin Qi, Tetsu Nakamura, Subodh K. Saha, Retish Senan, Yuhei Takaya, Hailan Wang, Hongliang Zhang, Mei Zhao, Hara Prasad Nayak, Qiuyu Chen, Jinming Feng, Michael A. Brunke, Tianyi Fan, Songyou Hong, Paulo Nobre, Daniele Peano, Yi Qin, Frederic Vitart, Shaocheng Xie, Yanling Zhan, Daniel Klocke, Ruby Leung, Xin Li, Michael Ek, Weidong Guo, Gianpaolo Balsamo, Qing Bao, Sin Chan Chou, Patricia de Rosnay, Yanluan Lin, Yuejian Zhu, Yun Qian, Ping Zhao, Jianping Tang, Xin-Zhong Liang, Jinkyu Hong, Duoying Ji, Zhenming Ji, Yuan Qiu,

Volume 62 · Number 4 · April 2024

Special Issue: Subseasonal-to-Seasonal predictability of extreme precipitation and land forcing

Guest Editors: Yongkang Xue · William K-M Lau

EDITORIAL

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Improved subseasonal-to-seasonal precipitation prediction

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LS4P Phase II (May-August 1998)

Major Objevtives:

(1). Case study for 1998: Cold Rocky Mountains and Warm Tibetan Plateau in Spring and drought in Southern Great Plains and June flood in Yangtze River, respectively.

(2) The LS4P research on other years and seasons, such as late summer and winter.

EVMWF-IFS Sensitivity Statistics (Preliminary)

	May 1998 2m T	June 1998 Precip mm/day			
	Tibetan Plateau		S. Yangtze Basin		
Obs. Anomaly	1.404		5.668		
Bias in CONTROL	-3.314		-5.281		
Sensitivity Experiments	Experiment minus CONTROL				
TP ∆t n=1	0.618		0.672		
TP ∆t n=3	0.717		1.759		

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Observed and Simulated Jun 1998 precipitation anomaly (mm/day) Over Yangtze River Basin and South Great Plains

	Obs anomaly	RM Cold LST/SUBT effect	TP Warm LST/SUBT effect	SST effect
Yangtze River Basin (26-31N;104-120 E)	3.26	1.11	1.44	1.04
South Great Plains (27-37N;107W-80W)	-1.45	-0.71	-0.61	-0.11

Note: The 1998 was a very strong El Niño Year. The SST has very strong impact in the tropical regions, such as Sahel, Central Africa, Amazon, and Central America at S2S scale (not listed in the table).

Summary

- 1). Observational data show a potential for the LST/SUBT to provide land memory at the S2S time scale, and a lag relationship between May T2m anomaly over the TP/RM and June precipitation anomaly downstream.
- 2). In ESM experiments, by properly producing observed May TP T2m anomaly, the S2S predictions over hot spot regions are improved. The consideration of the TP LST/SUBT effect has produced about 25%-50% of observed precipitation anomalies in most of 8 hotspot regions. For comparison, 6 regions with significant SST effects were identified in the 2003 case, explaining about 25-50% of precipitation anomalies over most of these regions.
- 3). The TP LST/SUBT influence is underscored by an observed out-of-phase oscillation between the TP and RM surface temperatures and a downstream TP-RM Circumglobal (TRC) wave train linking the TP to North America.
- 4). LS4P II focuses on TM effect and RM-TP Interactions. We welcome more groups to participate in.